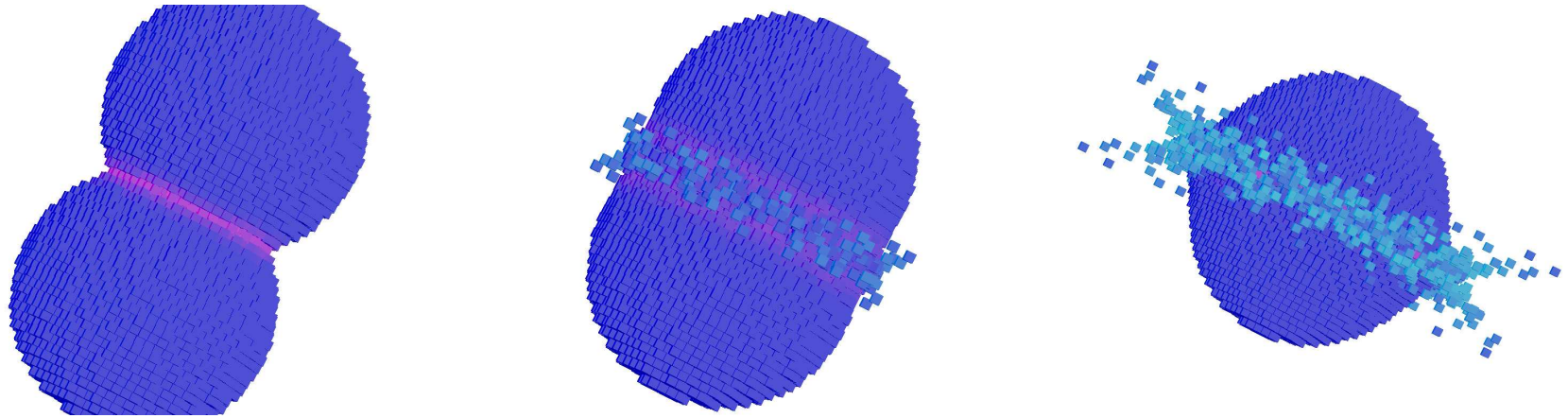


Simulations of Collisions Between Pre-Planetesimals



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Question for collisional coagulation from cm to km:

What is dominating in collisional outcome, accretion or erosion?

Answer depends on several parameters:

- Size and shape of target and projectile.
- Impact velocities (velocity dispersion due to gas drag in the disc and due to collisional history).
- Impact parameter / angle.
- Material composition and structure of target and projectile.

⇒ Model for 2-body collisions of pre-planetesimals needed (e.g. for global evolution simulations of the planetesimal disc).

Modelling collisions

Two regimes have to be distinguished:

- **Large objects** ($R \gtrsim 1 \text{ km}$). Main influence: self-gravity.
→ **Gravity dominated regime**
- **Small objects** ($R \lesssim \text{few } 100 \text{ m}$). Main influence: material properties.
→ **Strength dominated regime**

Objectives:

- What is the collisional outcome? **Growth** or **erosion**?
- **Distribution of size** and shape of possible fragments.
- **Velocity distribution** of possible fragments.

No laboratory **experiments** for sizes $R \gtrsim 1 \text{ m}$ possible.
⇒ **Numerical modelling** necessary.

Currently favoured numerical method:

Smooth Particle Hydrodynamics (SPH)

- Originally developed to simulate hydrodynamic problems in astrophysics. Later extended to elasto-plasto-dyn.
 - 3D mesh-free Lagrangian particle method
- ⇒ The continuum of the flow is represented by **interacting “particles”**, which move like point masses according to the **Lagrangian** form of the **equation of motion**, but represent all physical properties of that fluid part.
- Due to mesh-free **Lagrangian** nature:
- ⇒ natural reference frame for representing **deformations** and **fragmentation**.
- Due to particle nature: Easy to include **self-gravity**.

Collisions of brittle solid bodies (I)

First astrophysical applications:

- Collisional evolution of asteroids and asteroid families.
 - Gravitationally dominated regime.
 - Impact velocities higher (\sim km/s) than assumed for pre-planetesimal collisions.

⇒ Transition to strength regime at $R \sim 250 \pm 150$ m

(e.g. Love & Ahrens 1996).

⇒ More likely shattering impacts with gravitationally reaccumulated fragments than dispersing impacts.

Collisions of brittle solid bodies (II)

More general study: Benz & Asphaug 1999

- Collisions of **basalt** and of **ice spheres**,
 - $R \sim 1 \text{ cm} \cdots 100 \text{ km}$,
 - impact velocities of $\gtrsim 1 \text{ km/s}$.

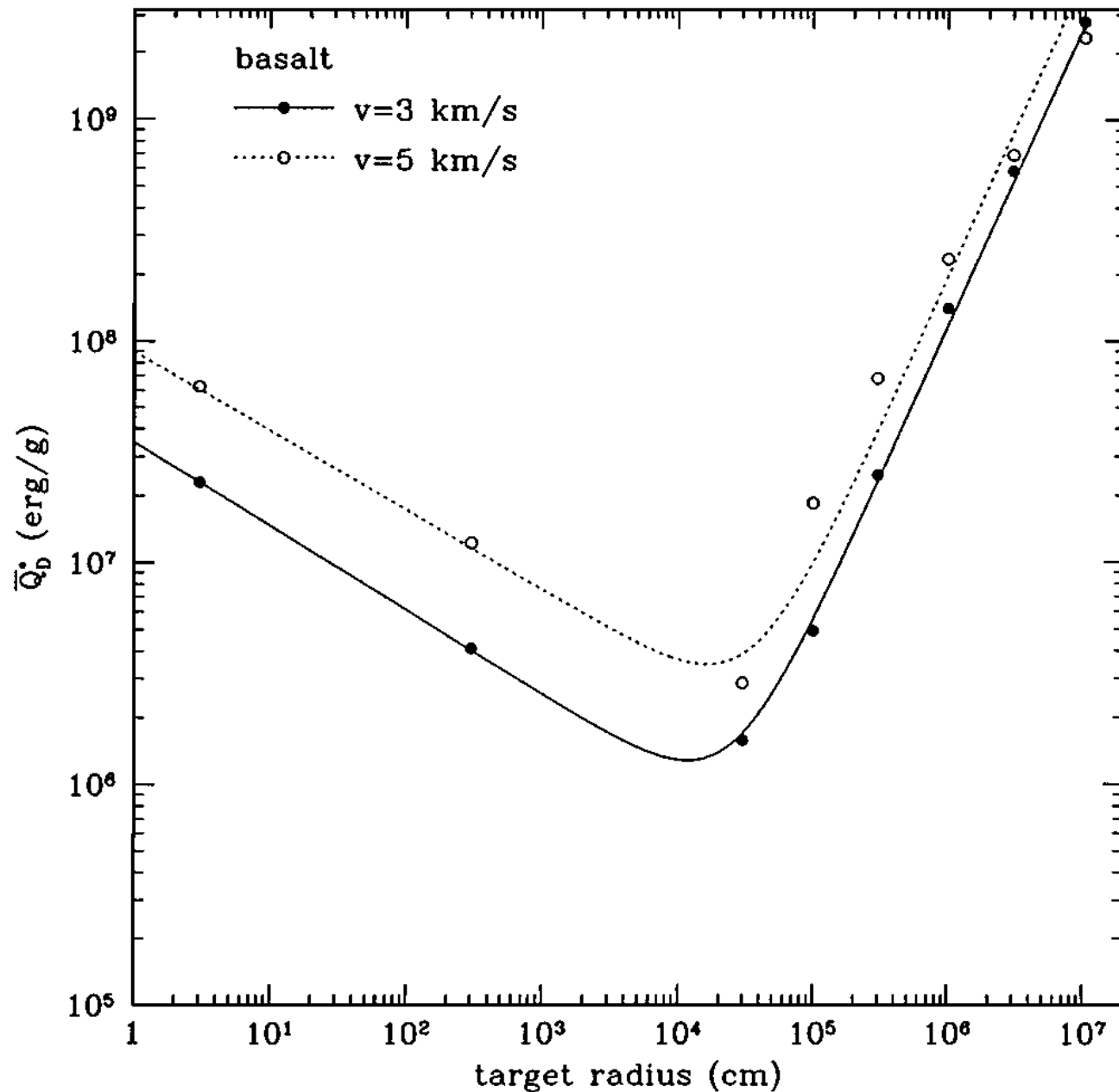
⇒ Similar conclusion: remnants of targets with $R \gtrsim 1 \text{ km}$ consist essentially of **shattered, gravitationally bound** fragments ⇒ **“rubble piles”**.

Characteristic quantity for impacts:

- **Catastrophic Disruption Threshold**

Specific incoming kinetic energy (kinetic energy in the collision per target mass), for which largest remaining object is half in size of target. (Note: remnant may well be fragmented!).

Catastrophic Disruption Threshold



⇒ weakest bodies at
 $R \sim 200$ m

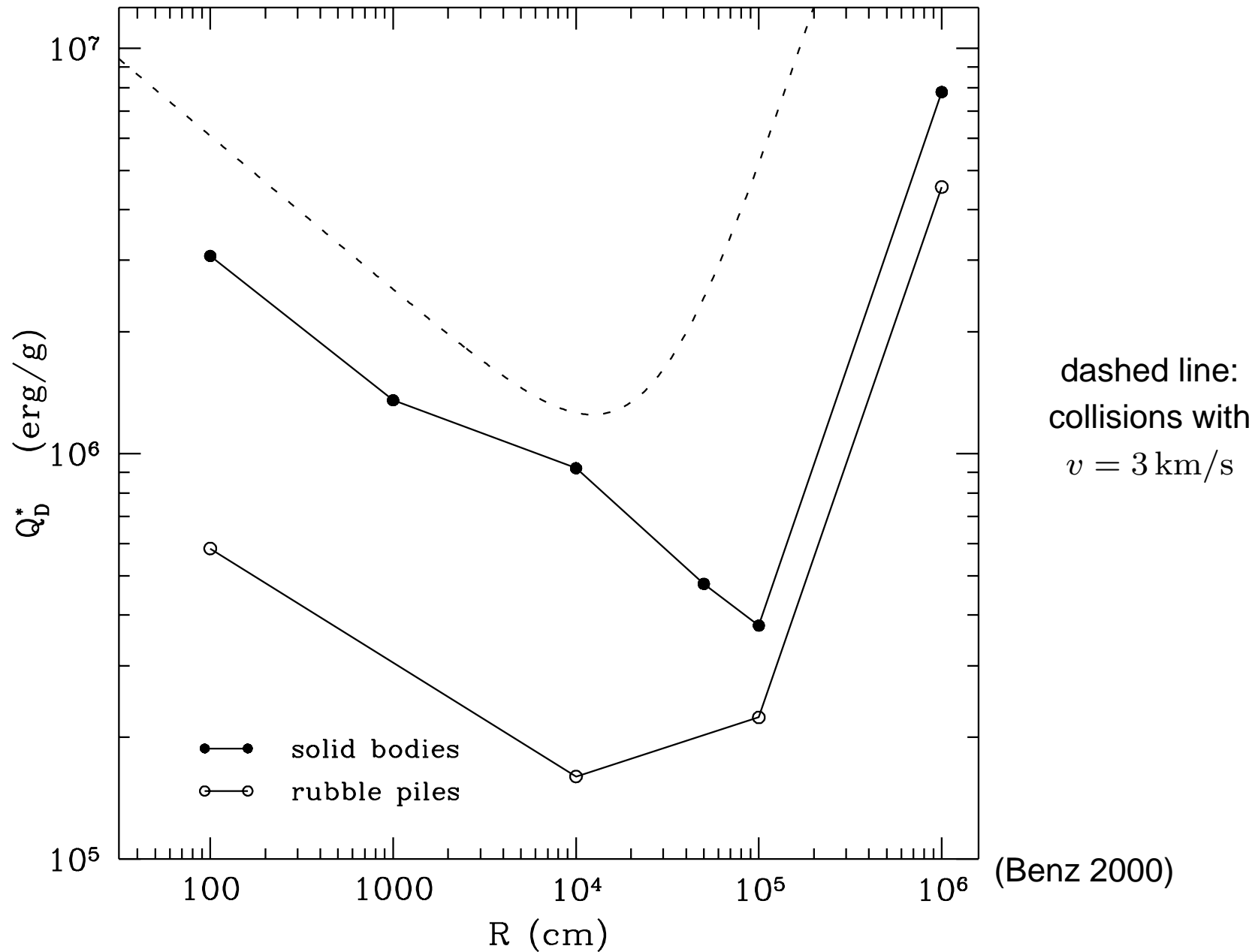
(Benz & Asphaug 99)

Low velocity collisions and rubble piles

Similar study but dedicated to the growth of planetesimals:
Benz 2000

- Collisions of **solid spheres** and of **rubble piles**,
 - basaltic material,
 - $R \sim 1 \text{ m} \dots 10 \text{ km}$,
 - collision velocities of $5 \dots 40 \text{ m/s}$.
- ⇒ Low velocity collisions **more vulnerable** to disruption.
- ⇒ Rubble piles considerably **weaker** than solid bodies.
- Reason: more **efficient momentum transfer** upon impact.

Catastrophic Disruption Threshold



⇒ Weakest objects: in the $R \sim 100$ m regime.

More problematic for pre-planetesimal growth:

Metre-sized bodies

- **Fastest radial drift** due to gas drag

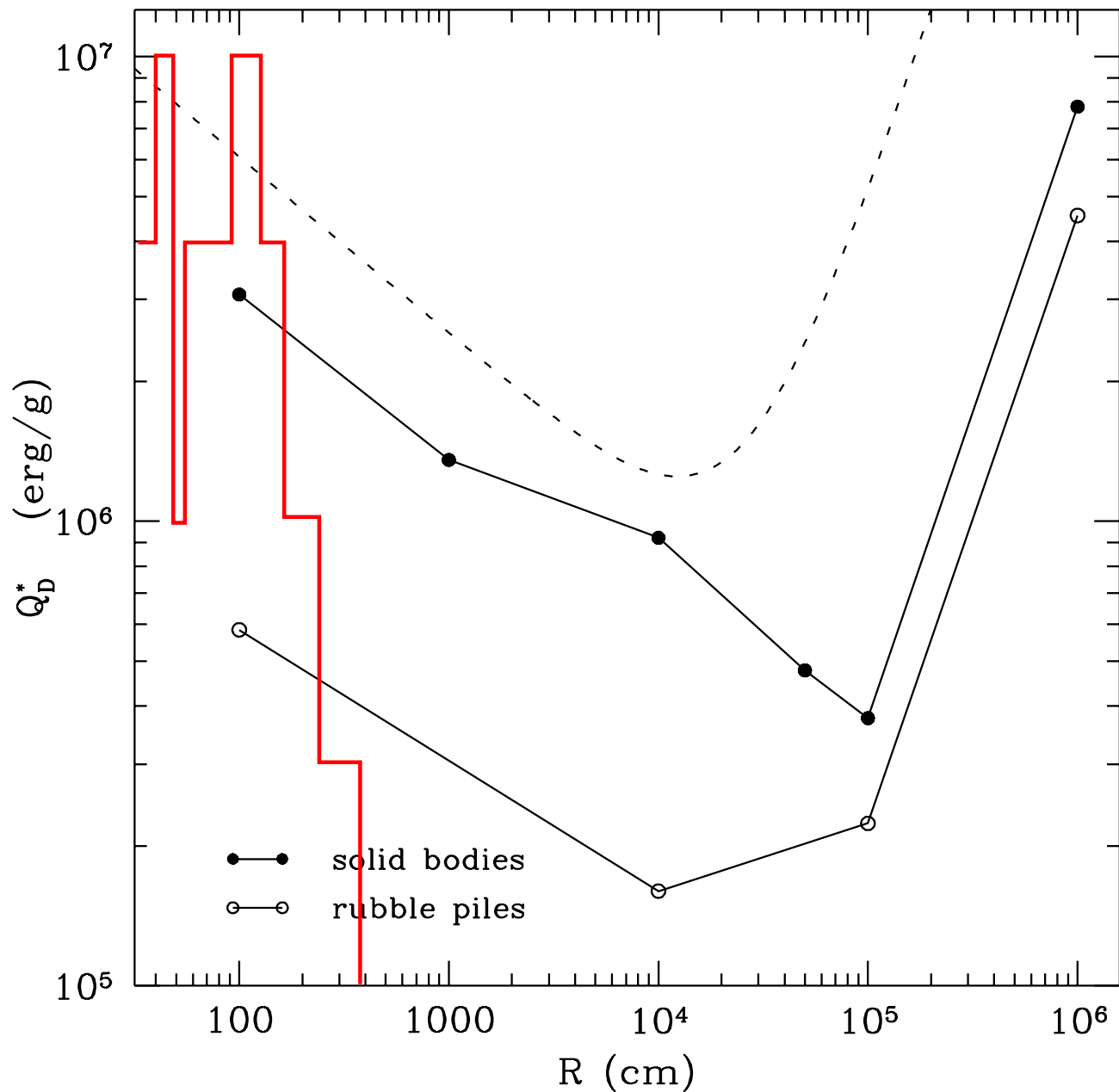
⇒ largest velocity dispersion and relative velocities
($\sim 10 - \lesssim 100$ m/s, depending on shape)

⇒ well **above** critical **disruption threshold**

⇒ easy **destruction** and **dispersion** upon collisions

⇒ **Bottleneck** in growth of pre-planetesimals?

Specific kinetic collisional energy due to gas drag compared with Critical Disruption Threshold

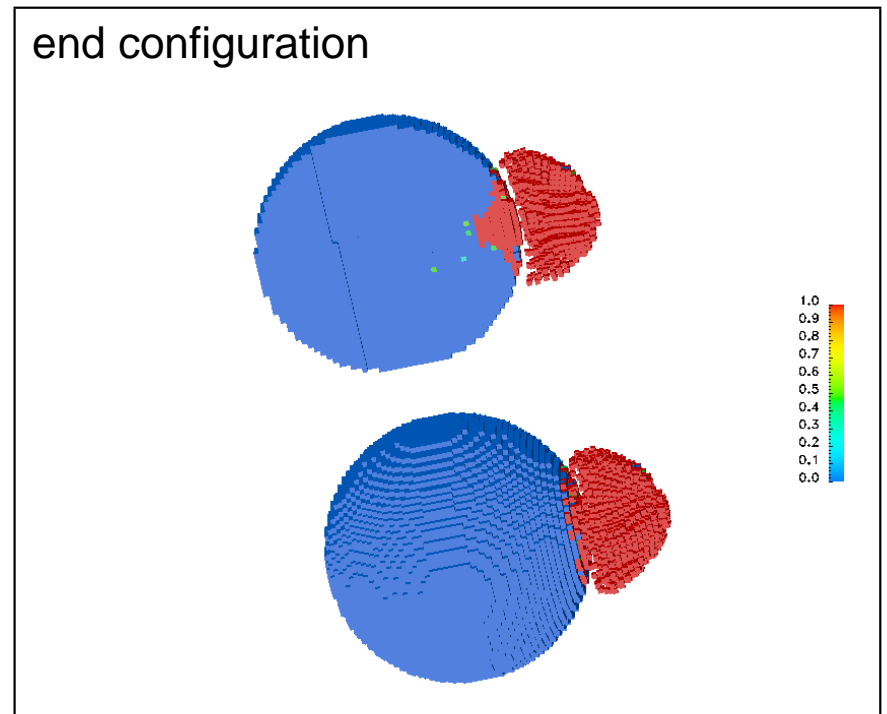
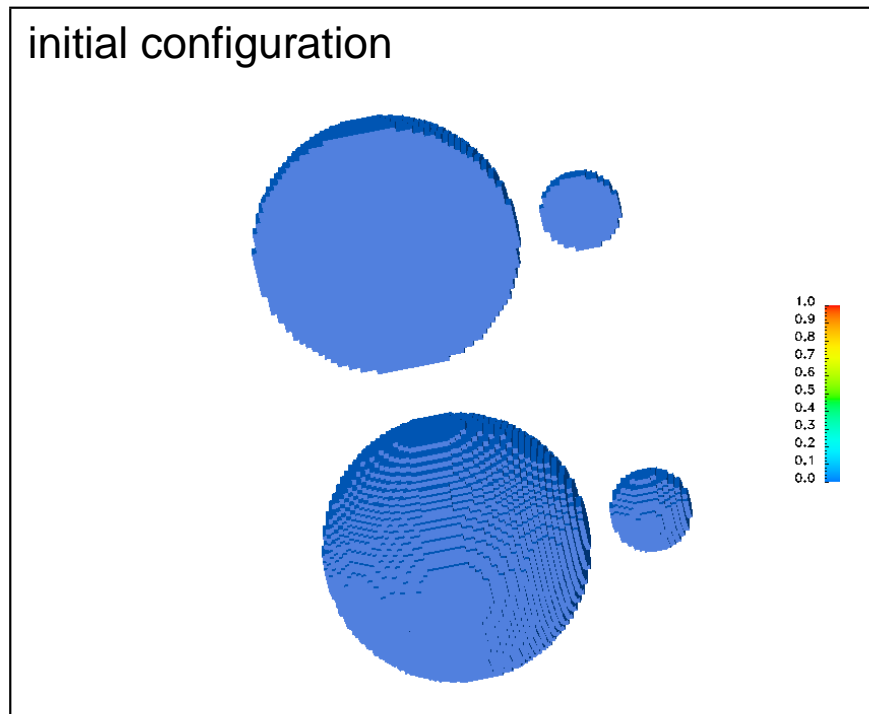


(red line: estimated from Fig. 7 of Benz 2000 for equal sized targets)

Collisions of porous (m -sized) bodies

- So far: simulations of collisions of solid bodies
 - consisting of **rocky materials**,
 - consisting of rocky **rubble piles**.
- ⇒ Results: **Erosion** or **fragmentation**, **no net growth**.
- However, pre-planetesimals may consist of **porous agglomerates** with differing material properties
(strongly indicated, e.g., by low density of asteroids and comets, by lab experiments of dust growth (Blum, Wurm), and theoretical simulations (Dominik, Tanaka)).
- ⇒ Next step: SPH simulations to study collisions of porous bodies.
- ⇒ **Porosity model** in SPH: material parameters depend on filling factor of density (Sirono 2004) (– see C. Dominik’s talk).
 - Problem: material parameters fitted from experimental data, but usually unknown.

Collision with small impactor – solid rocky material



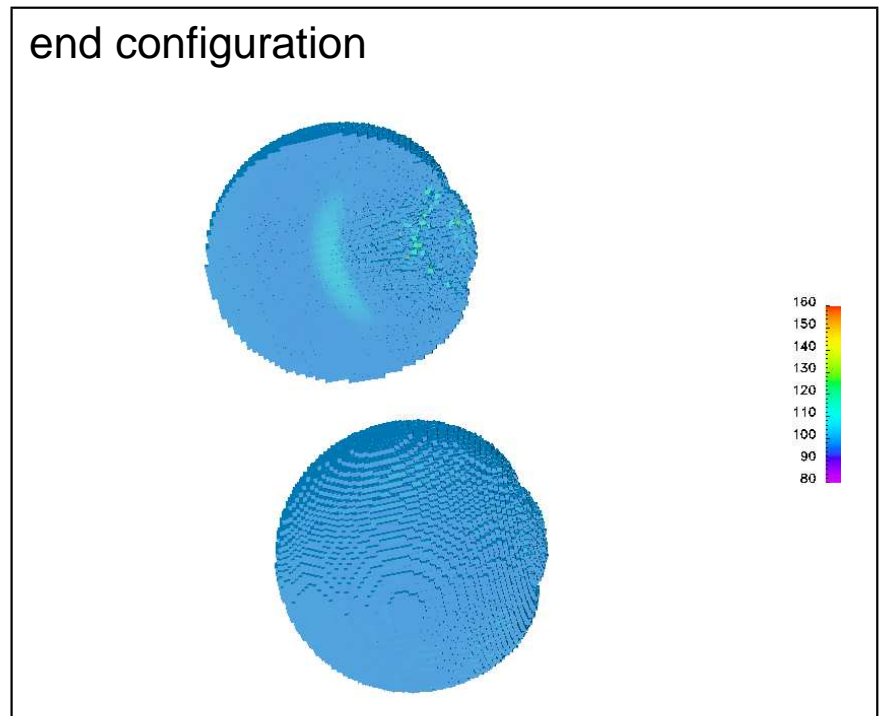
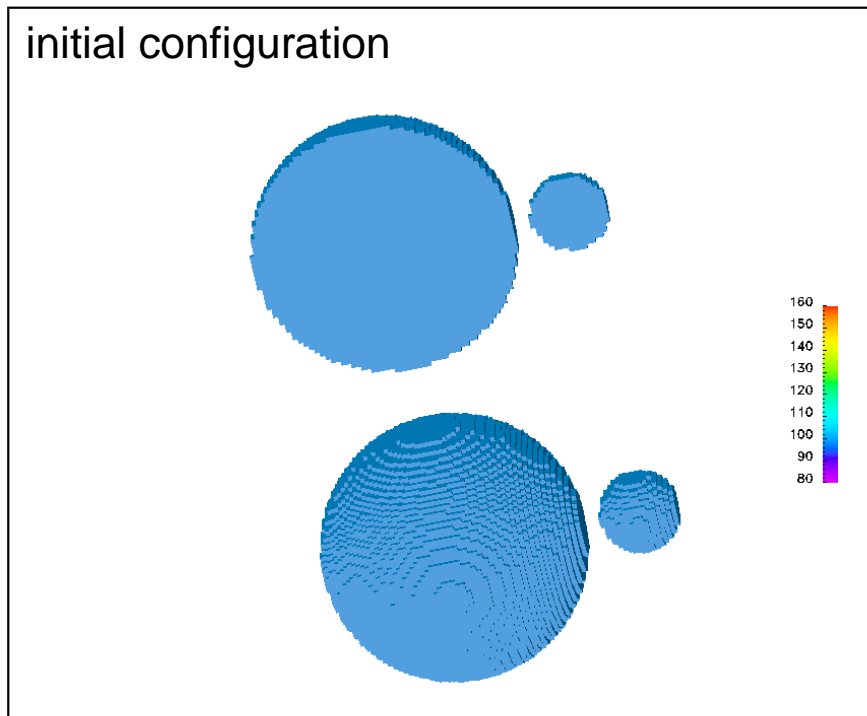
Colour-coded: damage

Target-radius: 1 m, Impactor-radius: 1/3 m, Initial density: 3 g/cm³,

Relative velocity: 20 m/s

(Movie)

Collision with small impactor – porous material



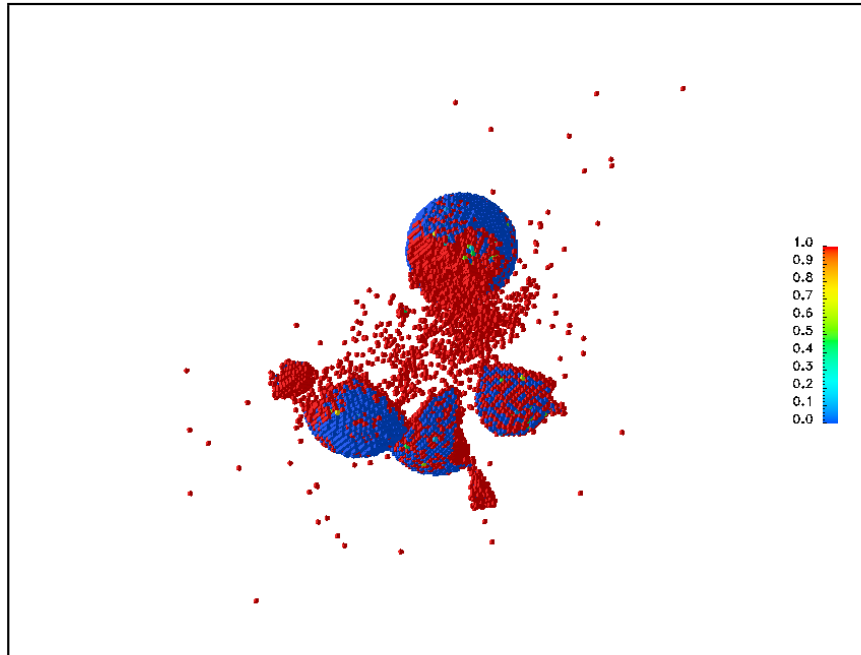
Colour-coded: density

Target-radius: 1 m, Impactor-radius: 1/3 m, Initial density: 0.1 g/cm³, Porous filling: 0.1,

Relative velocity: 20 m/s

(Movie)

Collision with equal size targets – oblique impact



Brittle material,
Target and impactor radius: 1 m,
Initial density: 3 g/cm^3 ,
Impact velocity: 20 m/s.

(Movie)



Porous material, filling factor: 0.1,
Target and impactor radius: 1 m,
Initial density: 0.1 g/cm^3 ,
Impact velocity: 20 m/s.

(Movie)

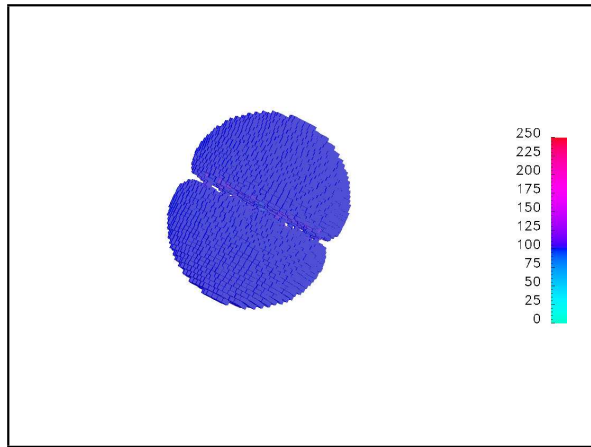
Influence of material parameters

Porous material, filling factor: 0.1, equal size targets (radius: 1 m), direct impact with 10 m/s.

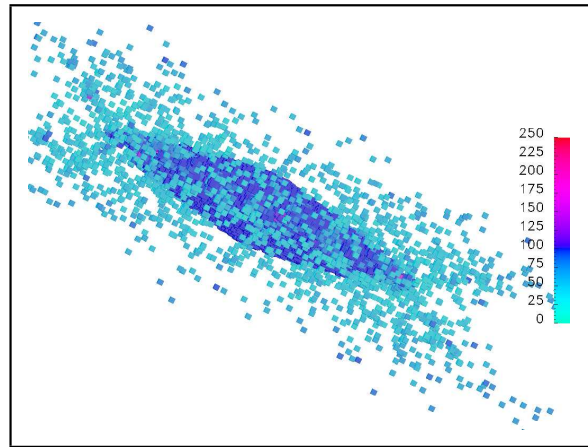
Compressive strength: $\Sigma(\rho) = \Sigma_0(\rho/\rho_0)^6$

Tensile strength: $T(\rho) = T_0(\rho/\rho_0)^5$

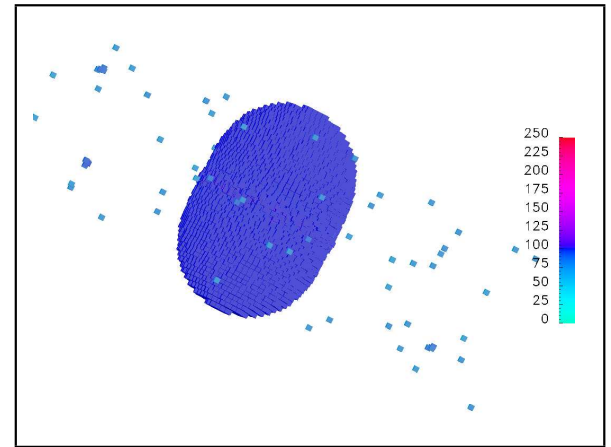
$\Sigma_0 = 6 \cdot 10^2, T_0 = -6 \cdot 10^3$



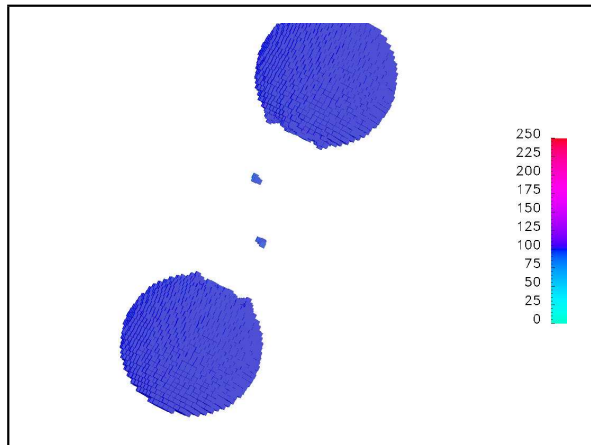
$\Sigma_0 = 6 \cdot 10^2, T_0 = -6 \cdot 10^2$



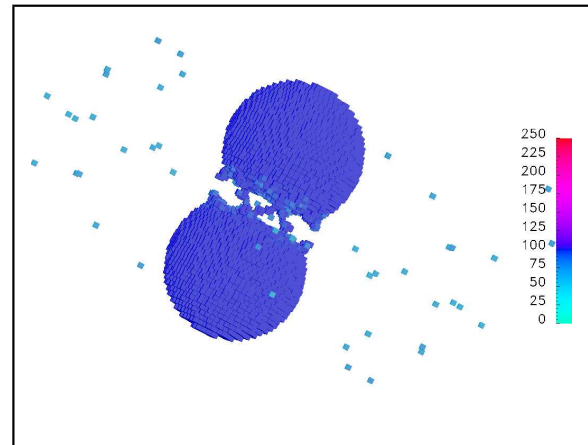
$\Sigma_0 = 6 \cdot 10^3, T_0 = -6 \cdot 10^3$



$\Sigma_0 = 6 \cdot 10^4, T_0 = -6 \cdot 10^4$



$\Sigma_0 = 6 \cdot 10^4, T_0 = -6 \cdot 10^3$



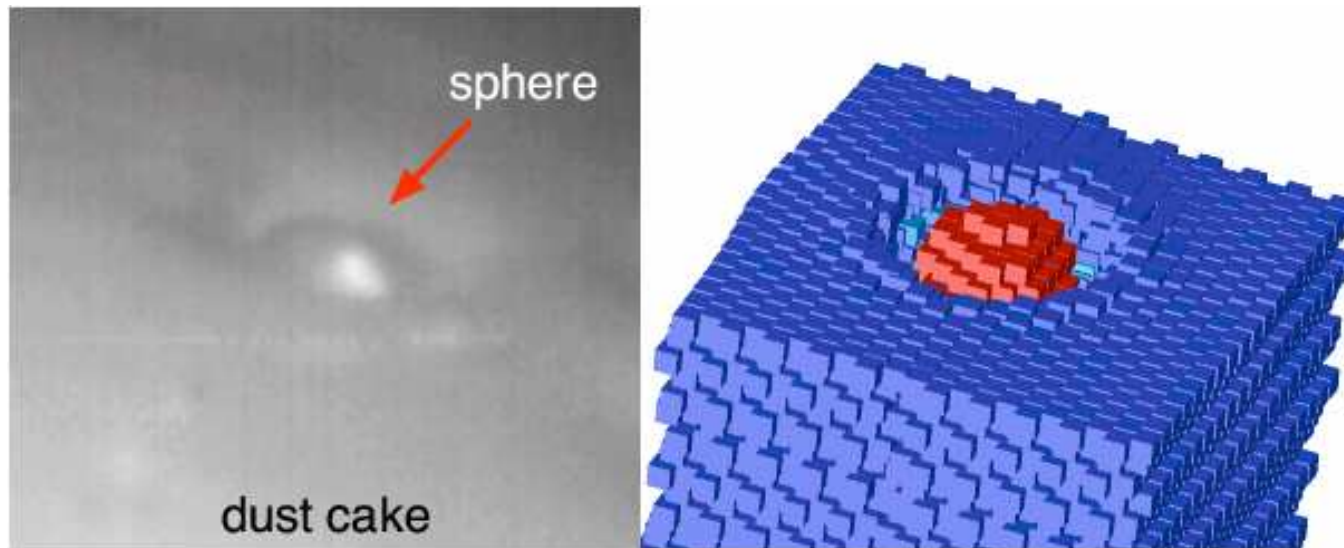
(All coefficients are in Pa)

Conclusion

1. For porous pre-planetesimals, growth is possible even by incorporating objects larger than grains.
2. Distinctly different collisional outcome for porous and rocky material, though indications for a general effect of erosion and fragmentation upon equal size collisions.
3. Collisions of metre-sized bodies still seem to be a barrier in growth of pre-planetesimals, however this is highly material dependent.

Outlook

- More **quantitative evaluation** of collisional outcome.
 - Methodical **parameter studies**.
(→ Catastrophic Disruption Threshold, etc.)
 - More **realistic material parameters**.
 - Improvements of **damage model** and **porous model**.
- ⇒ **Calibration** with laboratory **experiments**.



Blum et al.

Many Thanks

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